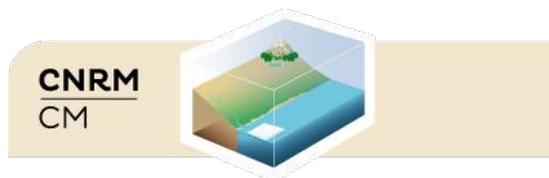


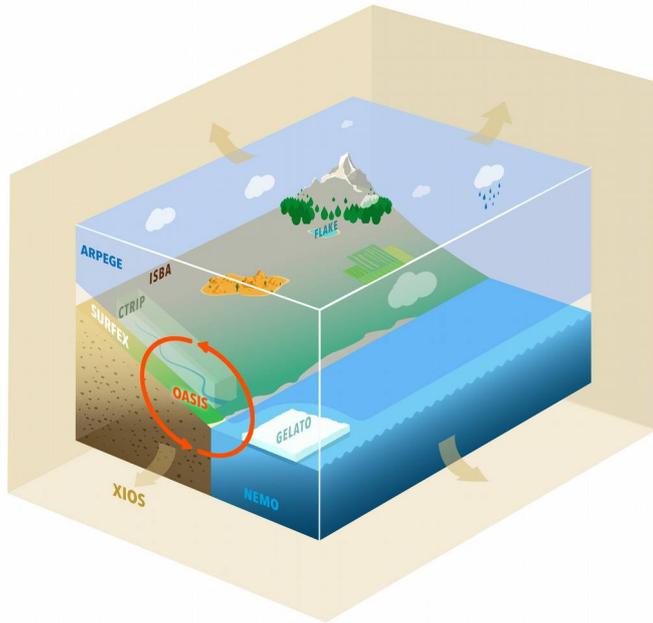
# Improving the river outflow coupling in a global climate model



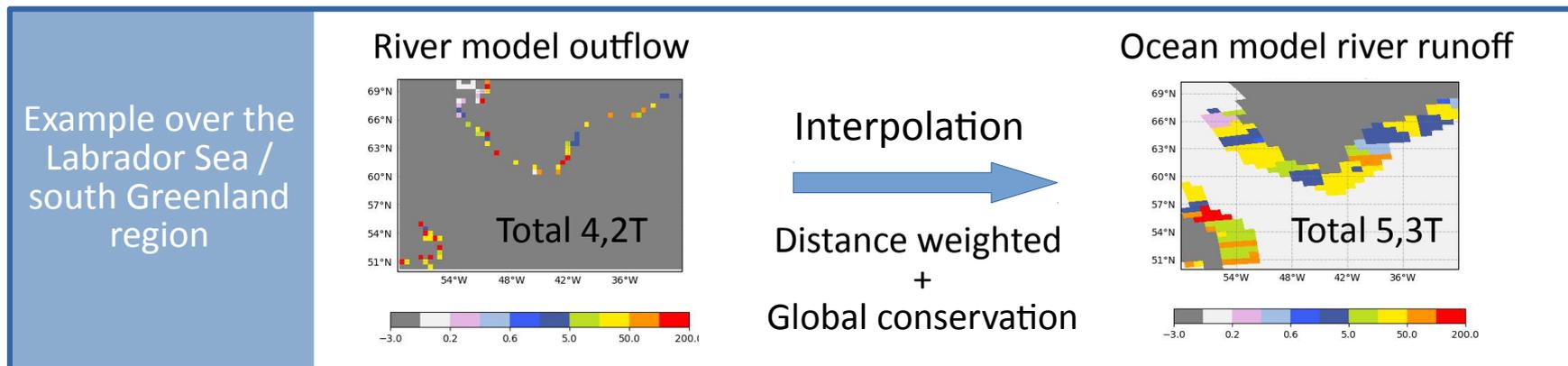
A. Voltaire [aurore.voldoire@meteo.fr](mailto:aurore.voldoire@meteo.fr)

# Background / Problematic

- CNRM-CM6-1 : the CMIP6 version of climate model developed by CNRM and Cerfacs (*Voldoire et al., JAMES, 2019*)
- Coupled components : Atmosphere-Land-River-Ocean-Sealce
- The coupling is done through OASIS3-MCT (*Craig et al., GMD, 2017*).
  - Coupling interpolation between components has been designed to be the simplest as possible to ease any resolution change in model components (ie use existing OASIS interpolation types)
  - BUT: there was no appropriate interpolation method for river outflow to ocean coupling, we used what was available
    - the river outflow was conserved globally but not regionally



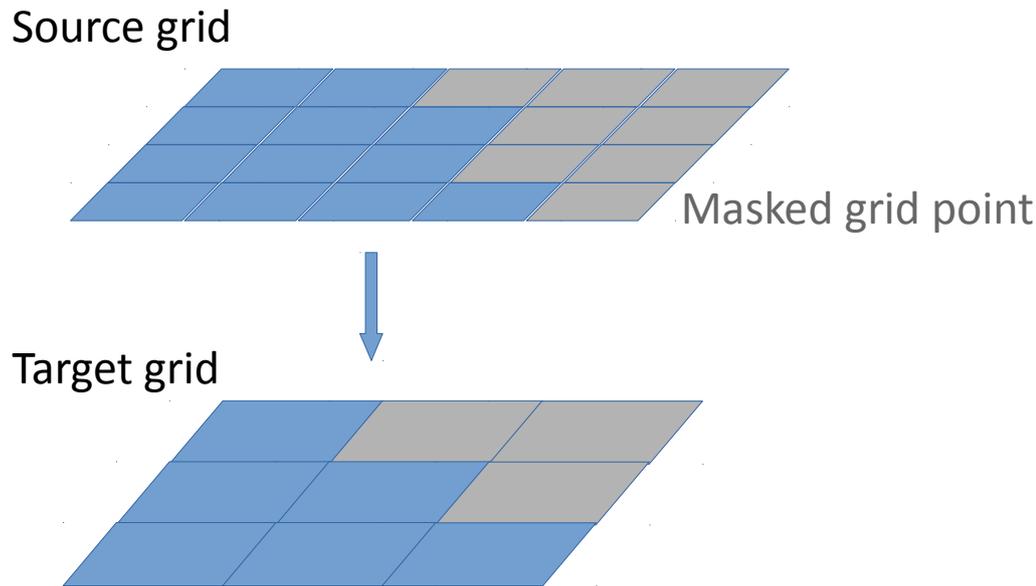
CNRM-CM



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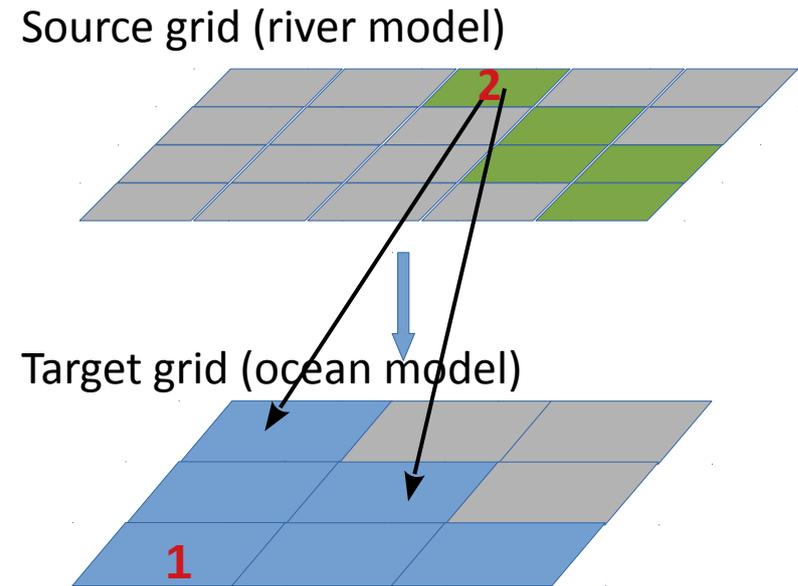
# Why existing interpolation methods in OASIS are not appropriate for river outflow interpolation to ocean?

General case



In general, the objective of the interpolation is to get a value on each unmasked target grid point  
This is the general point of view in OASIS

River outflow to ocean case



In the river outflow case,

- unmasked target grid points may not receive any flow (case grid point **1**)
- the source grid points are not geographically “over” the target grid points (case grid point **2**)

# New interpolation method

- For each source grid point, look for the closest target grid point
  - = the oasis nearest neighbour interpolation from the target grid to the source grid
  - = take the reverse interpolation links
    - x take the weights file for an Ocean → River nearest neighbour interpolation
    - x calculate new weights to conserve flux locally, for a given ocean grid point :

$$F_{ocean} = \sum_i ww_i F_{river}(i) \quad \text{where} \quad ww_i = \frac{area_{river}(i)}{area_{ocean}}$$

- NB:
  - x This can be extended to an interpolation with n-nearest neighbours.
  - x Easy to implement, links are already calculated in OASIS

# New interpolation method

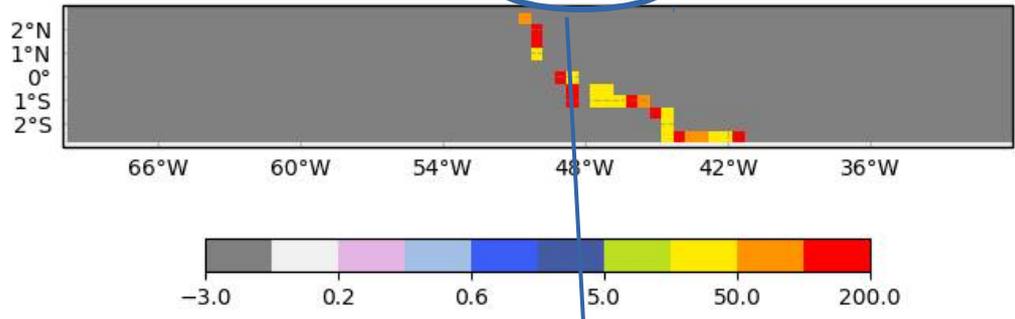
River  
outflow

Old  
interpolation  
method

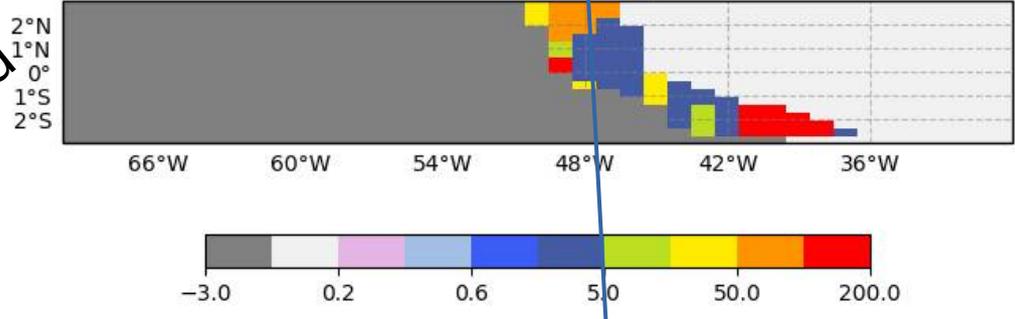
New  
interpolation  
method

Amazon

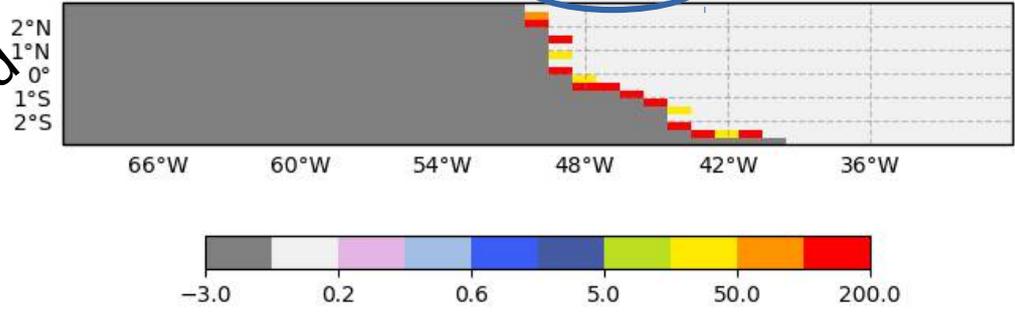
TRIP (98563 kg)



NEMO (35025 kg)

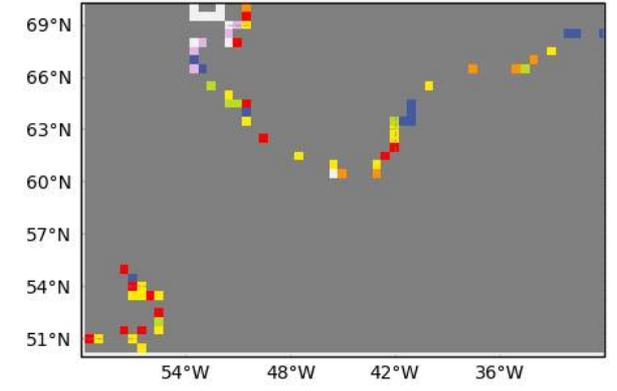


NEMO 1 (99934 kg)

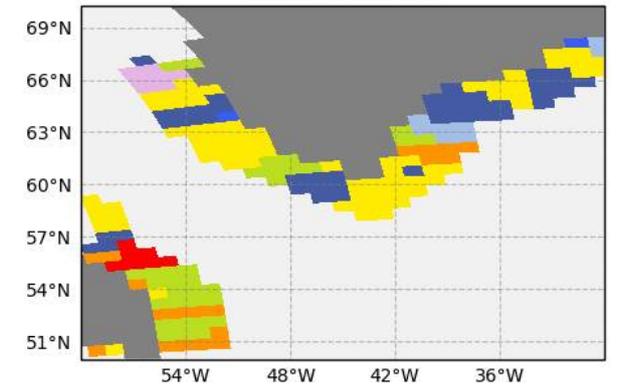


South-Greenland

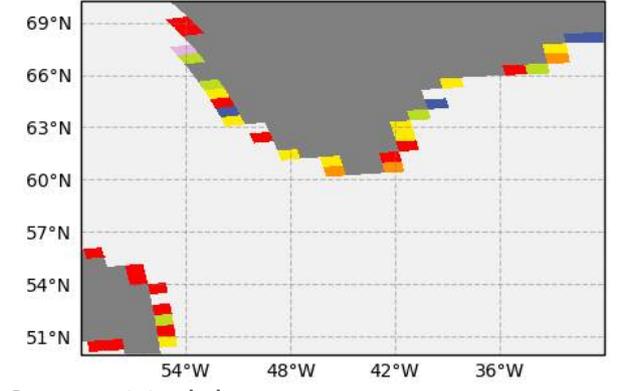
TRIP (4156 kg)



NEMO (5277 kg)



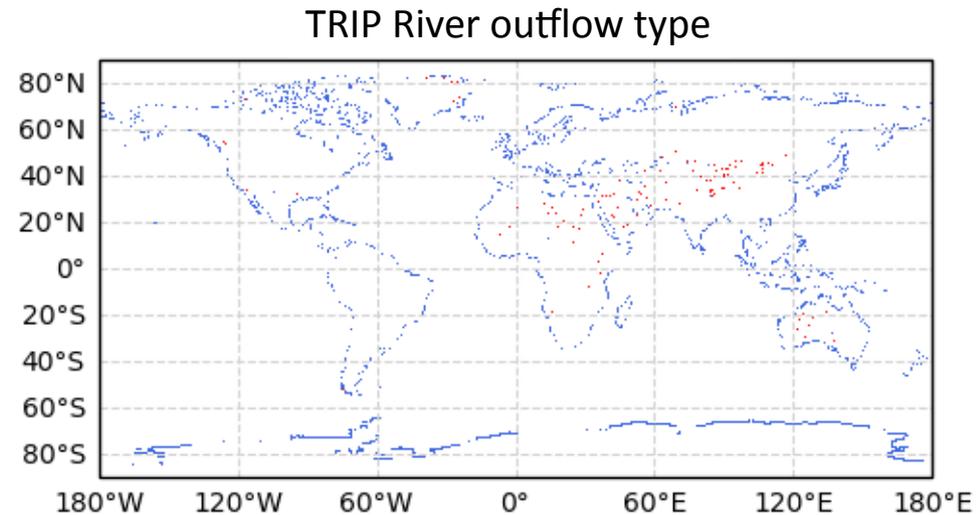
NEMO 1 (4292 kg)



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# Second improvement

- Not all TRIP river outflow are at the coast
- For **coastal** grid points, the nearest ocean point is close  
→ ok



- For **inland outflow**, the nearest ocean grid point is far  
→ not physical to spread water on one distant ocean grid point  
→ must take these river outflow into account otherwise the water mass budget is not closed in the model
- Solution proposed : spread the runoff from inland outflow over the whole ocean surface = small uniform quantity

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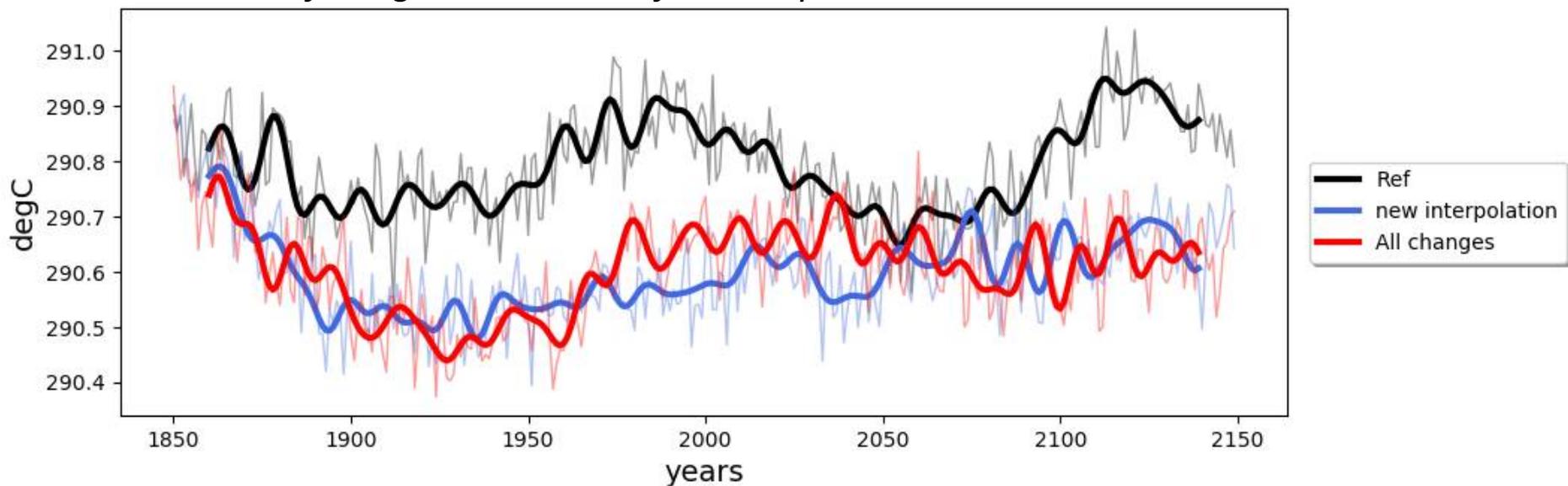
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# Sensitivity experiments performed

- **Ref** : reference experiment is the CMIP6 piControl simulation (forcings fixed to their preindustrial values)
- **New interpolation** : new interpolation method using **1** neighbour
- **All changes**, all recent developments gathered :
  - New interpolation method using **1** neighbour
  - Separate treatment for inland outflow
  - Land-sea mask of ocean and atmospheric model in better agreement

250 years of  
integration

*Evolution of the global mean surface temperature*



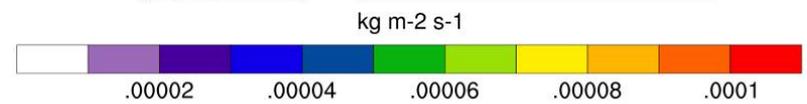
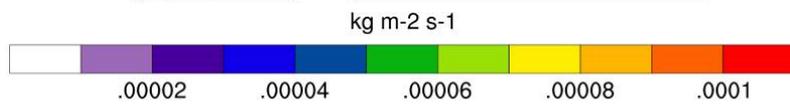
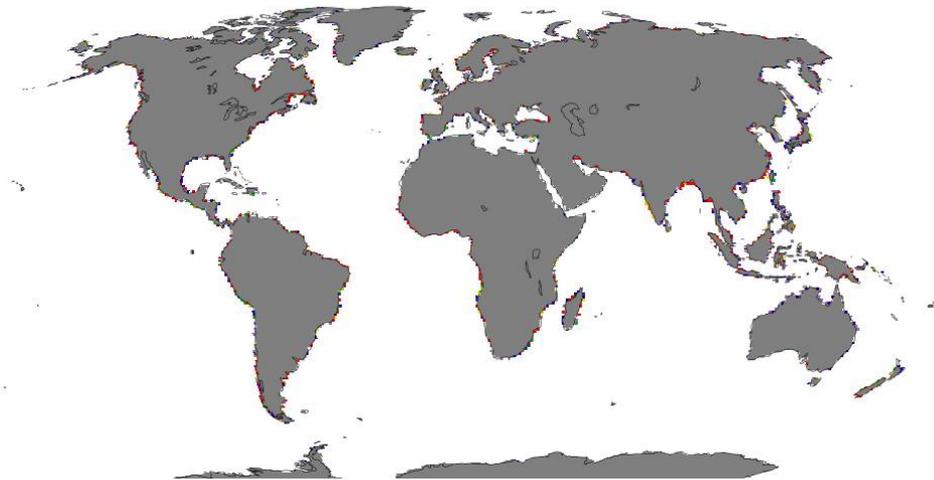
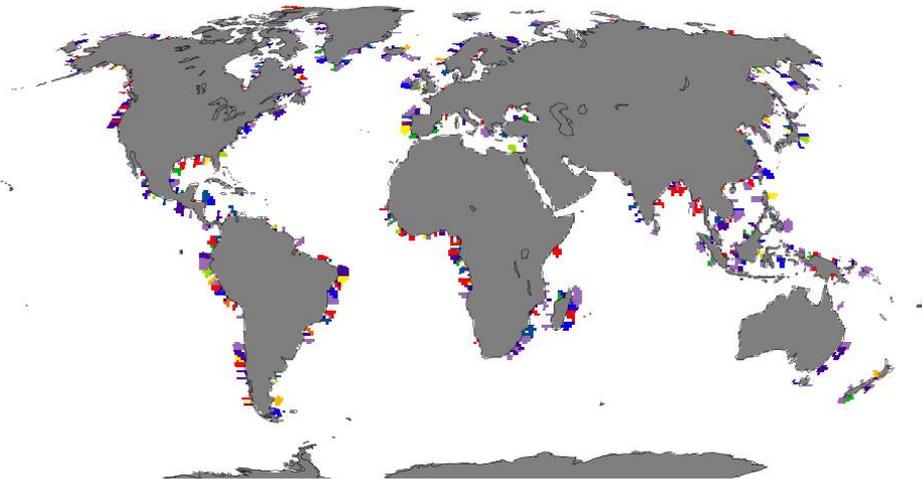
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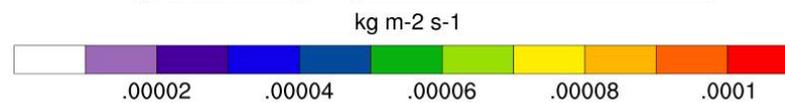
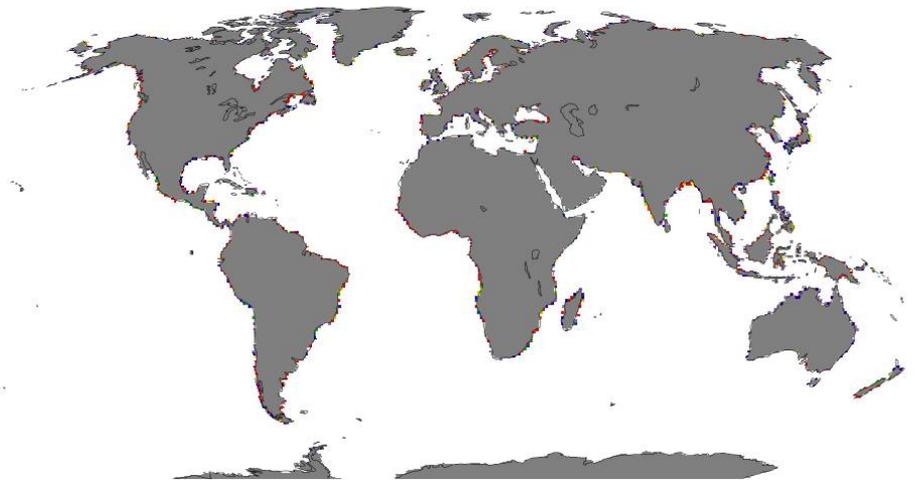
# Runoffs in the ocean (moy 1900-1949)

Ref

New interpolation

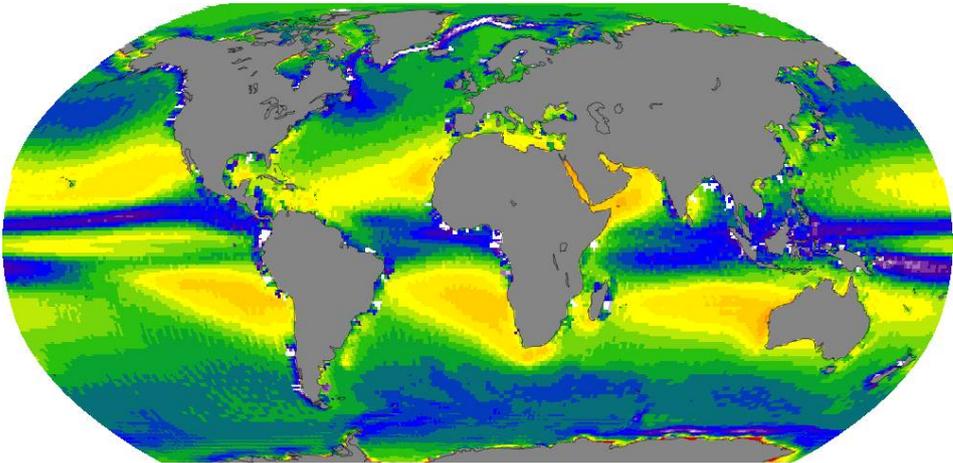


All changes



# Net water flux from the ocean (moy 1900-1949)

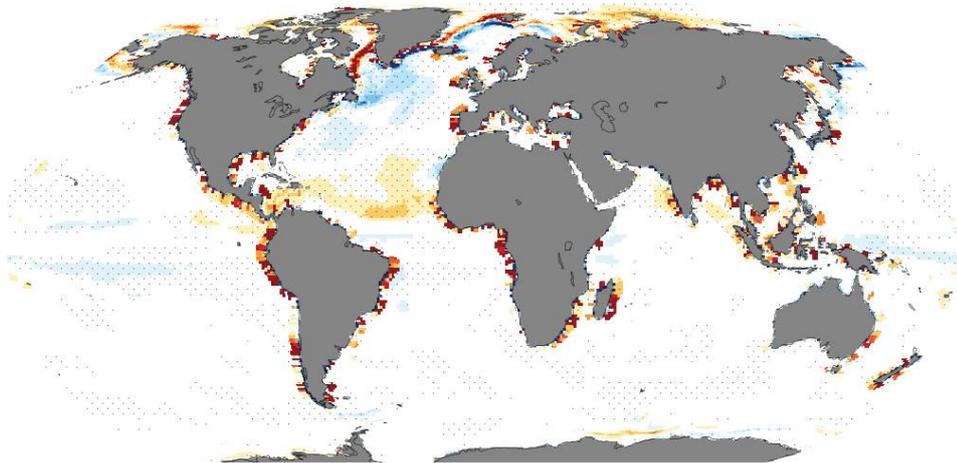
Ref



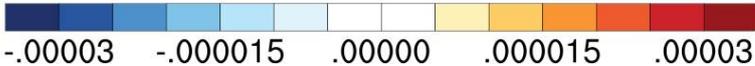
kg m<sup>-2</sup> s<sup>-1</sup>



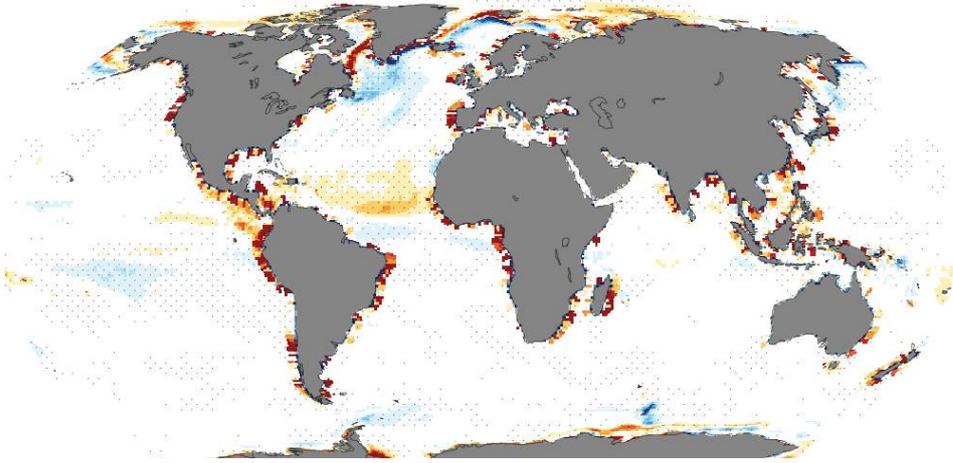
New interpolation - Ref



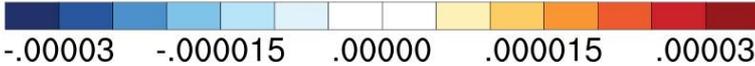
kg m<sup>-2</sup> s<sup>-1</sup>



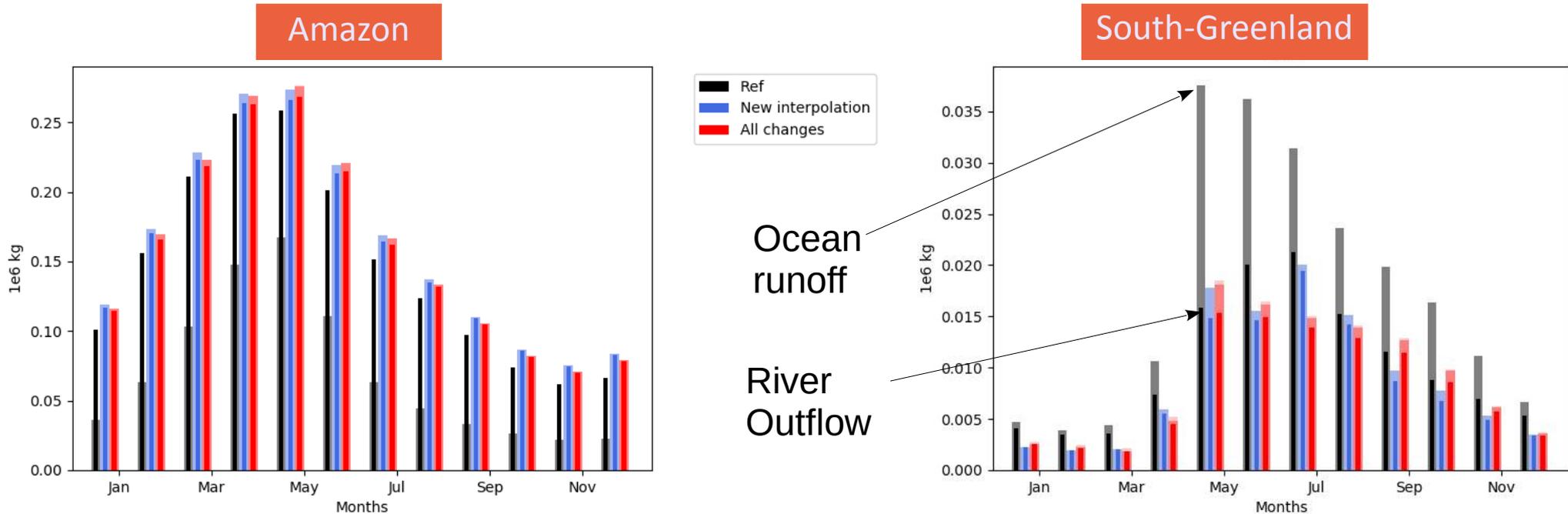
All changes - Ref



kg m<sup>-2</sup> s<sup>-1</sup>



# Outflow regional budget



*Mean annual cycle of river runoffs averaged over years 1900-1949 for each simulation and for 2 regions. For each simulation, the dark color bar indicate the river outflow and is superposed on the light color bar which represents the runoff received over the ocean.*

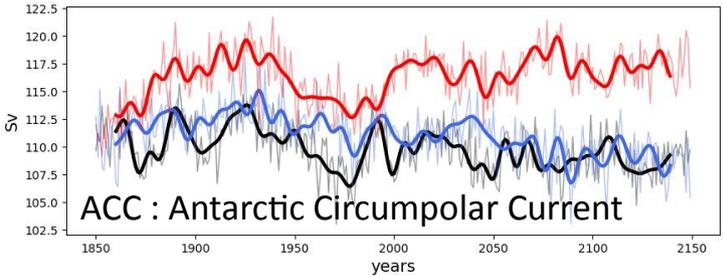
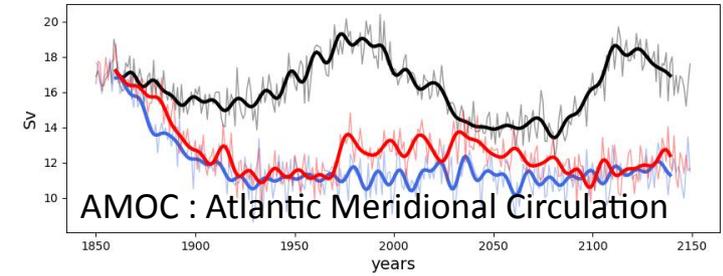
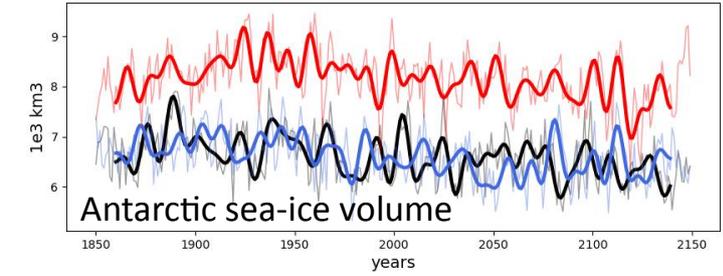
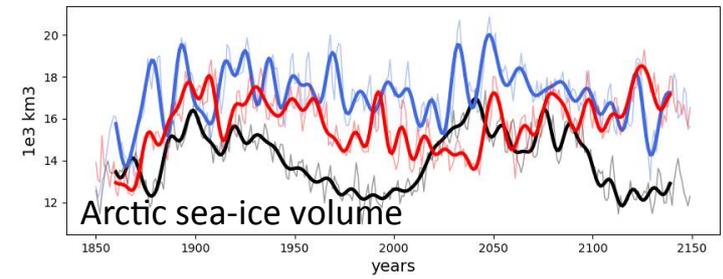
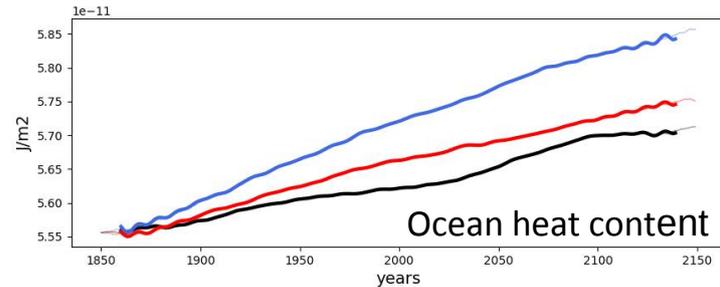
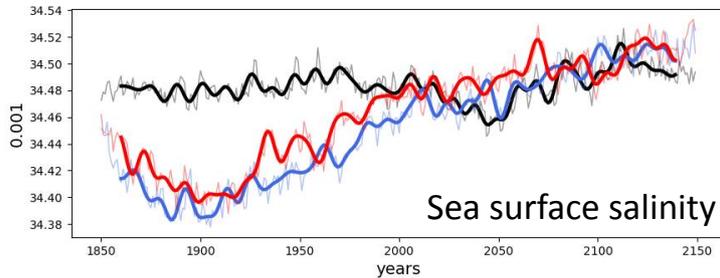
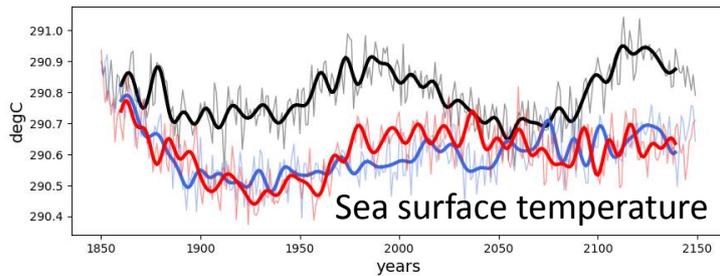
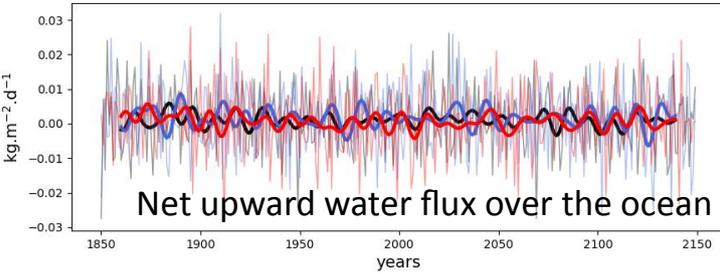
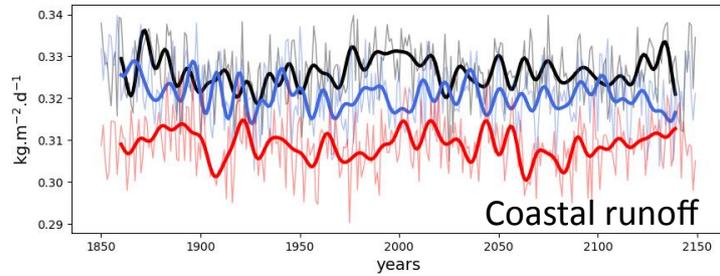
- x Confirm the regional conservation whatever domain is considered

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# Global impacts

Global annual mean (thin lines) and high-pass filtered time-series (thick lines)

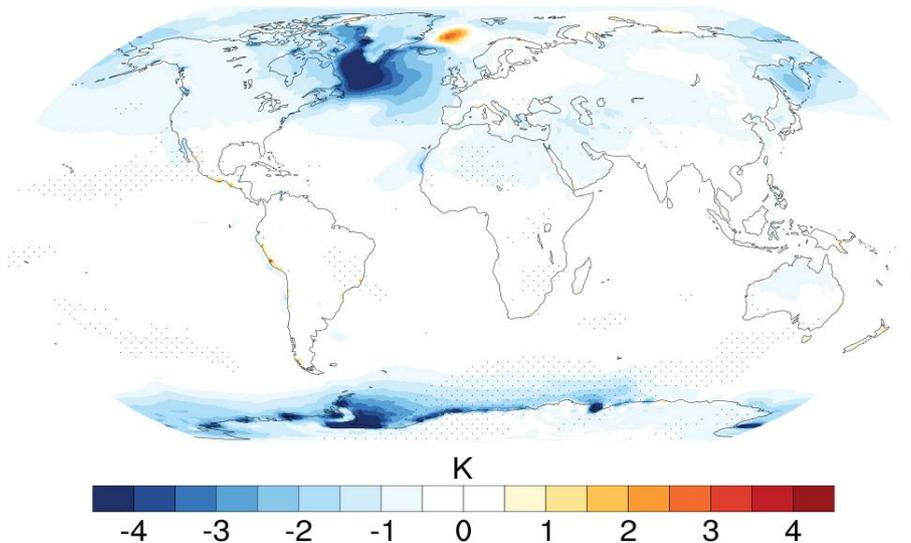


- × Strong perturbation of the mean climate
- × Takes 150 years for surface salinity to stabilize
- × Increase in Arctic sea-ice for both experiments
- × Increase in Antarctic sea-ice for the « All changes » experiment.
- × Change in sea-ice impacts the large-scale ocean circulation, ie reduction in AMOC and increase in ACC

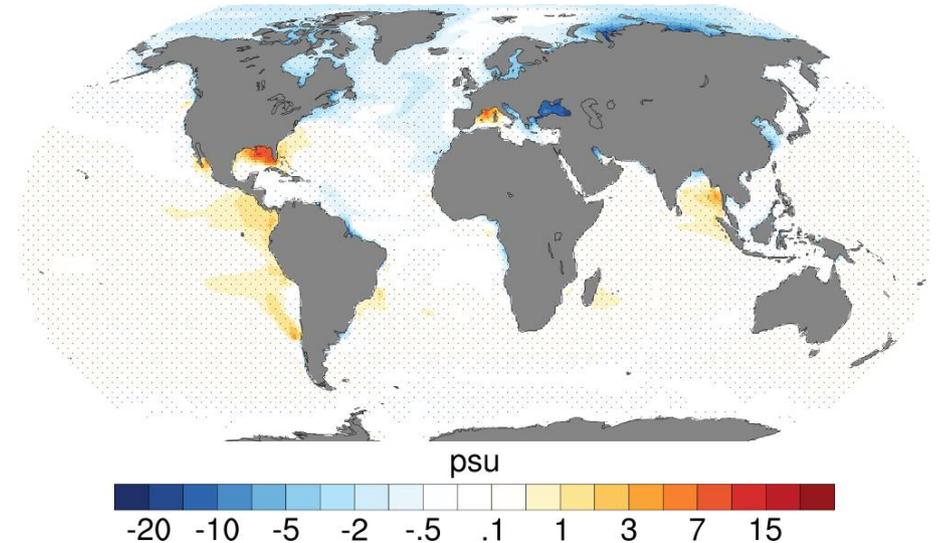
# Regional impacts

Mean difference **All Changes** – Ref  
averaged over the last 50 years

Near surface temperature



Ocean surface salinity



- × Strong impacts on surface salinity that spread far from the coast
- × Strong reduction in surface temperature over high latitudes → Reveals strong sea-ice changes

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# Sea-ice impacts - volume

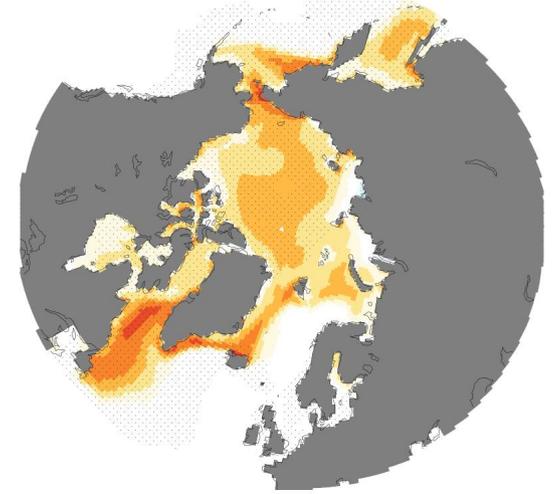
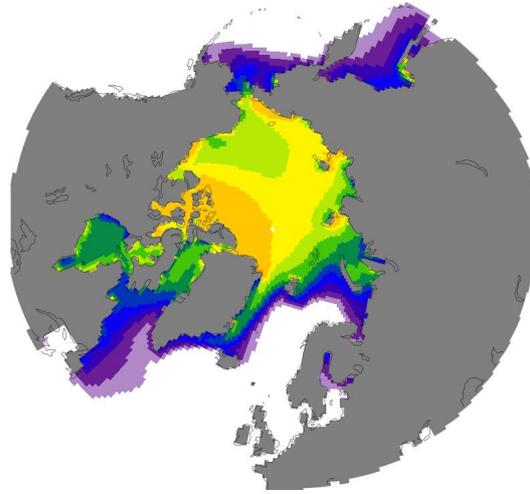
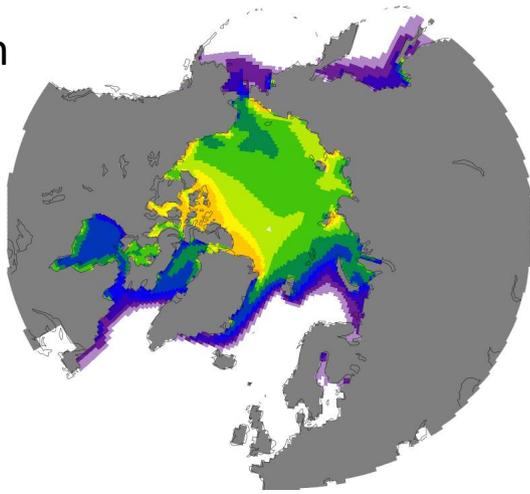
*Averaged over the last 50 years*

**Ref**

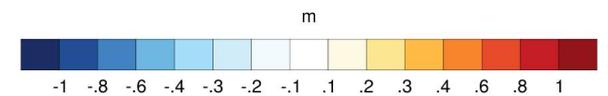
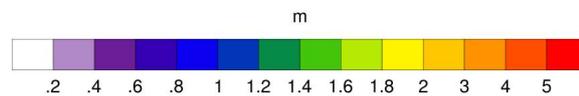
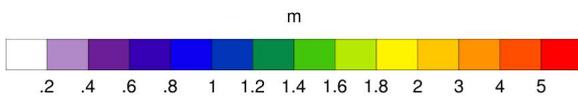
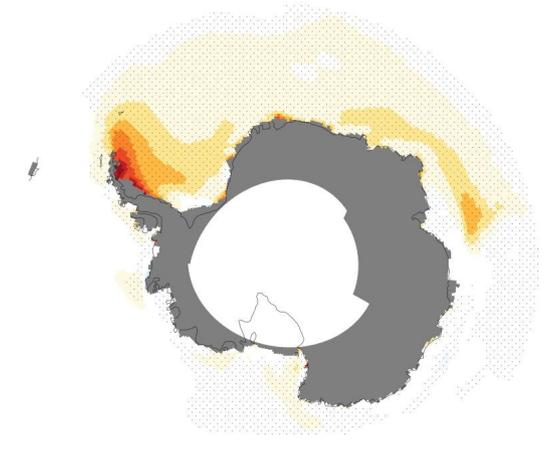
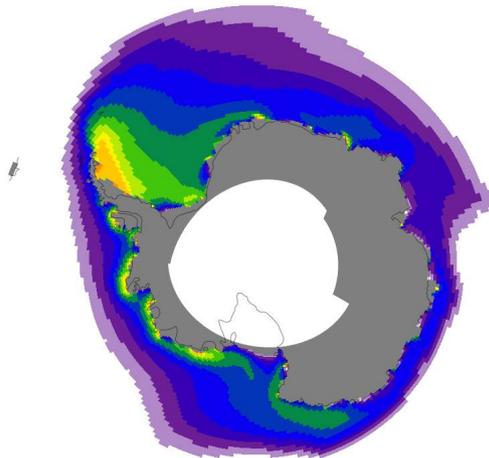
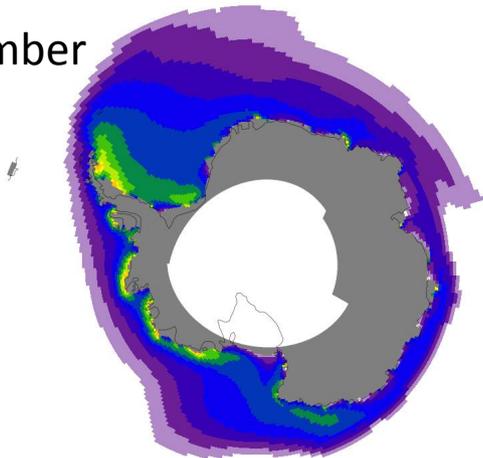
**All changes**

**All changes - Ref**

March



September



# Conclusions

- Propose a generic method to interpolate river outflow to ocean
  - based on OASIS nearest neighbours from the reverse interpolation
  - just need to re-calculate weights

This new method **ensure water conservation locally**

- The impact on the model mean climate is important on
  - salinity over high latitudes and closed sea
  - sea-ice cover and volume

Such changes impact mean ocean mass transport and circulation (AMOC, ACC)
- The new method will be directly available in OASIS in future versions.